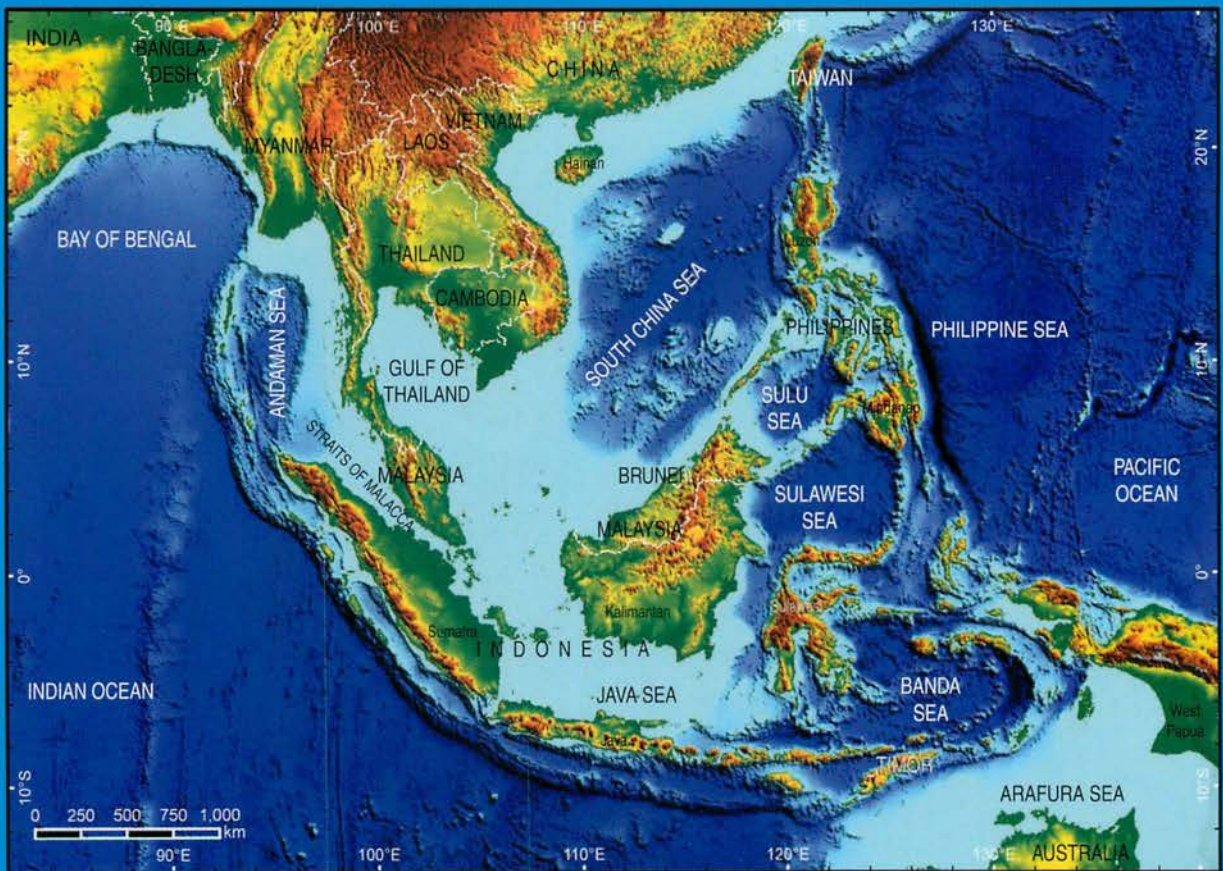


Geological Evolution of South-East Asia

Second Edition

CHARLES S. HUTCHISON



Geological Society of Malaysia
2007

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Second edition

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© Charles S. Hutchison 1989
First published by Oxford University Press 1989
This edition published with the permission of Oxford University Press 1996

ISBN 978-983-99102-5-4

*Printed in Malaysia by
Art Printing Works Sdn. Bhd.*

This book is dedicated to the former professors at the University of Malaya. It is my privilege to have collaborated with Professors C. S. Pichamuthu, T. H. F. Klompé, N. S. Haile, K. F. G. Hosking and P. H. Stauffer. Their teaching and publications laid the foundations for our present understanding of the geology of this complex region. I also salute D. J. Gobbett for having the foresight to establish the Geological Society of Malaysia and Professor Robert Hall for his ongoing fascination with this region.

Preface to this edition

The original edition of this book was published by Oxford University Press in 1989 as number 13 of the *Oxford monographs on geology and geophysics*. It was originally planned by me to include a section on economic deposits. However, the editor in Oxford instructed that this should be more comprehensively covered in a subsequent book. It appeared under the title 'South-East Asian oil, gas, coal and mineral deposits' in 1996 as number 36 of the *Oxford monographs on geology and geophysics*. Number 36 is still available from Oxford because it has been included in their new 'print when required' system. Number 13, unfortunately preceded this new system, and it is no longer available from Oxford.

With the permission of Oxford University Press, the 'Geological Evolution of South-East Asia' was reprinted in Kuala Lumpur, without any revision, and published by the Geological Society of Malaysia in 1996. This blue paper-back edition became well

known throughout the region of South-east Asia. Unfortunately the stock has become depleted in 2007.

It is too important a book to let go. A thorough revision would be beyond my time constraints. I have revised only those sections that have become absolutely necessary because of important new data.

Petaling Jaya
2007

CSH

ACKNOWLEDGEMENTS

I am grateful to the former President, Professor Lee Chai Peng, for support and considerable help in bringing this new edition into publication, and to Mrs Anna Lee for her friendly assistance. Many thanks are due to Dr. Ng Tham Fatt for designing the book cover.

Preface to the original Oxford book

My involvement in South-east Asian geology began in 1957 with systematic studies on the granite and gabbro-norite of Singapore Island. Gradually, these studies were extended throughout the Malay Peninsula and into Thailand and South China. My interests broadened to the volcanic rocks, stratigraphy, structure, and economic deposits of the region.

After completing the *Geology of the Malay Peninsula* with D. J. Gobbett in 1973, it became my desire to write a book on the greater region, for no regional geology had ever been written on South-east Asia. Warren Hamilton completed his landmark *Tectonics of the Indonesian region* in 1979, paving the way for the present book.

Professor P. H. Stauffer and I made initial plans for a book under the title *South-east Asia: geology and environment*. Unfortunately, the compilation of this book had to be abandoned in 1983 when Professor Stauffer left the region.

An enquiry from the Oxford University Press for a book under my sole authorship conveniently arrived in 1984, and was readily formalized.

This is not a standard regional geology, and the selection of the title allowed me to emphasize tectonic models, but at the same time to document the stratigraphy and igneous events. I have always taught my students that every aspect of geology involves interpretation. From thin sections to hand specimens, to maps, and thence to regions there is an increasing order of interpretation. Although it has greatly benefited from and has been influenced by numerous discussions with colleagues, the mistakes are mine alone.

This book gives more of a bird's-eye view of the region, for much field and laboratory work remains before a definitive account can be written. There remains great scope for high-quality palaeomagnetic research to refine the Phanerozoic plate motions modelled in the following pages. The granites have been reasonably well dated, but there is still a lack of radiometric dates for volcanic and metamorphic formations in addition to granites in critical tectonic regions. Good structural geological analysis is also lacking, but lack of continuous outcrops will severely hamper progress.

The present volume links two major regions included in the present *Oxford Monographs on Geology and Geophysics Series*: No. 2 *Phanerozoic*

Earth history of Australia, edited by J. J. Veevers (1984), and No. 3 *The Geology of China*, by Yang *et al.* (1986).

The present volume will I hope appeal to regional stratigraphers and students of the great Tethyan province. The inclusion of Tertiary oil and gas basins may result in the use of this book by exploration companies. Unfortunately, mineral deposits have been excluded due to length constraints.

This volume is the outcome of courses I have taught to my students at the University of Malaya over the past 5 years. By the time it is published, I will have closed my own books and—with regret—have left the region after 30 enjoyable years of involvement in the geology of South-east Asia. It is my earnest hope, however, that this volume will remain open on the desks of my graduates and future generations of geology students to stimulate their better understanding of this complex and fascinating region.

Exeter
1987

C.S.H.

ACKNOWLEDGEMENTS

The compilation of this book has benefitted from collaboration and discussions with many colleagues. In particular, I would like to single out P. H. Stauffer, I. Metcalfe, K. R. Chakraborty, and Y. G. Gatinsky. Many of their ideas have been crystallized, albeit imperfectly by me, in the pages of this book. Ian Metcalfe kindly critically read the manuscript. Thanks are due to J. K. Blake for help in compiling Table 5.1.

S. Srinivass, Ching Yu Hay, and Roslin Ismail formed my sketches into the final illustrations. Jaafar Abdullah and Lee Kok Eng gave me photographic assistance. Rujiraporn Polchai helped in compiling the index. Zaimah Ahmad Saleh and Fauziah Hanim Saidi helped with the typing at various stages. My son Timothy helped in the final stages of the manuscript. I wish to thank the editors and staff of the Oxford University Press for their help in producing this book. A debt of gratitude is expressed to the University of Malaya for facilities to carry out my research and writing throughout the years.

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Introduction

1.1. GENERAL

The region described extends approximately 4330 km from latitude 26°N to 13°S, and approximately 5570 km from longitude 85°E to 135°E (Fig. 1.1). Its area, approximately 24 million km², is roughly comparable with that of Europe, and somewhat larger than that of Australia. Although geology should recognize no political boundaries, it is useful to be aware of the political subdivisions, for geological explorations have been and are largely restricted to individual countries. More often than not the international boundaries represent the limits of geological study, or changes in terminology.

The region includes all the Association of South-East Asian countries (ASEAN), namely Indonesia, Thailand, Malaysia, Singapore, and the Philippines. In addition, Myanmar and Indochina (Cambodia, Laos, and Vietnam) naturally fall into this region. For completeness, some reference will be made to Taiwan, the southern and eastern regions of China including Tibet and Hong Kong, the Assam area of India, Nepal, the Andaman and Nicobar Islands, and the northern parts of Australia and New Guinea.

About two-thirds of the region is covered by seas, in which there are thousands of islands, ranging in size from Sumatra and Borneo to those which are too small to show on the map. The interwoven network of seas, islands, and the peninsulas of mainland Asia and Australia have allowed a successful synthesis of modern oceanographic research with classical geological studies on land. Our understanding of South-east Asia has been significantly advanced as a result of international research co-ordinated and supported by the Co-ordinating Committee for Geoscience Programmes in East and Southeast Asia (CCOP), and the Intergovernmental Oceanographic Commission of UNESCO (IOC), as part of the International Decade of Ocean Exploration (IDOE). The programme came to be recognized as Studies in East Asian Tectonics and Resources (SEATAR), and a progress report was published some years ago (CCOP-IOC 1980).

The impressive amount of oceanographic work carried out in this SEATAR project has made possible the correlation of previously completed geological investigations on land with the submarine topography and its tectonic elements.

South-east Asia provides the world's most outstanding geological laboratory for the study and understanding of active plate tectonics. Deep-sea trenches, marginal seas, and island arcs abound. The region is seismically active and contains many of the world's most active volcanoes. The island arcs, and small intervening seas, are in the process of being compressed by the 8 cm per year northwards movement of cratonic Australia. The immediate effects of the arrival of Australia are to be seen in collision tectonics on the island of Timor and extinction of the volcanic arc to the north of it. As Australia continues to push northwards, the complex of island arcs and seas will be compressed further, eventually forming a major orogenic welt sutured between mainland Asia and Australia. By this time the seas will have been driven out, and the disparate islands and basins will be compressed into a Himalayan-style complex such as can now be seen between India and Tibet. The study and understanding of South-east Asia is therefore important in the development of the geological sciences, for this is a region in which a future mountain or orogenic system can be seen in the early stages of formation. Every orogen is unique, but a study of the sequence of events in the South-east Asian archipelagos and seas will help in the unraveling of older now completed orogens.

The region is also of interest in providing evidence of the relationships between tectonic evolution and economic deposits. Sedimentary basins lying between island arcs and occupying shallow shelf seas on continental margins are important for oil, gas, and coal. The volcanic island arcs are important for porphyry copper and precious metal deposits. Base metal deposits are also of some importance. Uplifted oceanic lithosphere contains chromite and nickel deposits. In the more continental parts of the region, there are spectacular granite-related tin and important tungsten

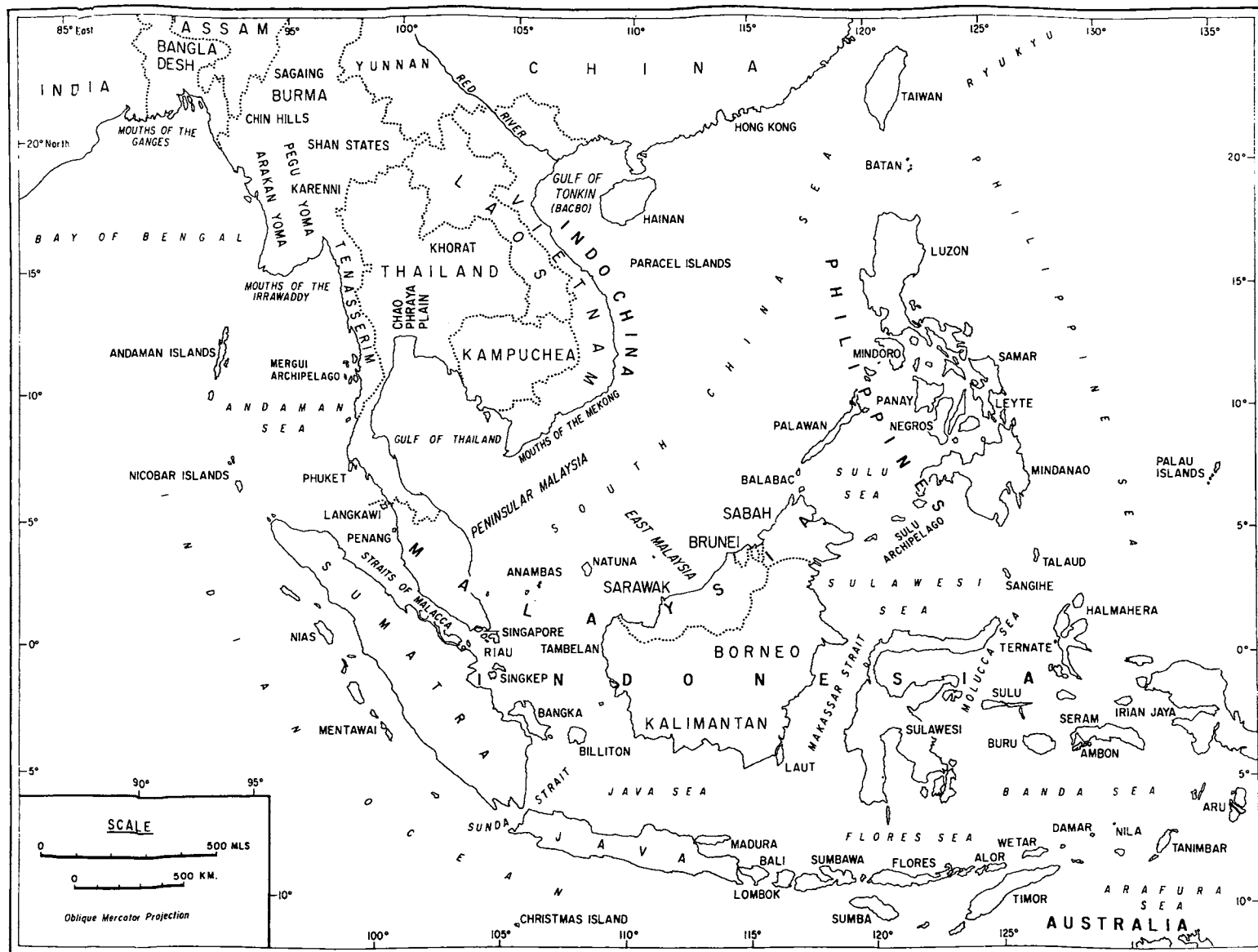


Fig. 1.1. Outline map of South-east Asia showing the countries and some of the place names. Burma is now known as Myanmar and Kampuchea as Cambodia.

deposits, and antimony and mercury are locally significant. Some of these deposits have been locally worked out.

Environmentally, the region may be divided into a stable cratonic core and an outer island arc system, which is seismically and volcanically active. Some major fault zones, such as the Red River, have high seismic risk. The volcanic arcs of Indonesia and the Philippines offer rich volcanic soils, but at the same time they are regions of high earthquake risk. The archipelagos which face the major oceans are susceptible to tsunami damage. Volcanic eruptions have frequent local effects, but they are more widely felt if the magnitude and frequency of the eruptions increase (Blong and Johnson 1986). Large-scale clearances of jungle and forest result in significant changes in micro-climate, and most certainly result in soil erosion. Urbanization, especially in flat coastal plains and inland valleys, has led to severe air and water pollution, and changes in the micro-climate. Withdrawal of groundwater at excessive rates has led to regional subsidence, which in the case of Bangkok is beginning to result in disastrous flooding and structural collapses (Natalaya and Rau 1981).

1.2. HISTORY OF GEOLOGICAL STUDIES

The majority of the South-east Asian countries benefited from their colonial eras in that established geological surveys, complete with experienced geologists and administrative infrastructure, were established in the territories. Only Thailand was spared the yoke of colonialism, with the drawback that systematic geological study and publication have accordingly been slow to develop.

Geological and mining activities were totally disrupted by the Japanese occupation of the whole region in the early 1940s and valuable records and manuscripts were lost for ever. Continuing insurgency in many South-east Asian countries has until recently prevented systematic geological surveys in parts of the region, and political disputes prevent a proper assessment of the oil potential of the major part of the South China Sea.

1.2.1. Thailand

The first useful geological description was by Hallet (1890). He was a civil engineer in search of a railway route from Moulmein to south China, and he traveled through north Thailand by elephant, describing the geological formations seen on the way.

The Royal Department of Mines and Geology was established in 1891, but little geological work was done. Its principal function was to grant tin mining leases. The second director made several trips through the country and his publications contain some descriptions of the mineral deposits (Smyth 1898).

An Anglo-French convention of 1896 secured the borders with Myanmar (then Burma) and Cambodia, and assured the continuing independence of Thailand. The southern borders with Malaysia (then called the Malay States) were long disputed, but were fixed by a treaty with Britain in 1899. In 1909 Thailand ceded to Britain its rights over the northern states of what is now Peninsular Malaysia.

The first critical geological investigation was made in 1912 by a professor from Uppsala University (Hogbom 1914). Prince Kamphaengbej, who was commissioner of the Royal State Railways, engaged Wallace M. Lee, an American geologist, to investigate the coal and oil resources of Thailand, in an attempt to find fuel for the locomotives. Lee made three investigations between 1921 and 1923, the first to the north, the second to the peninsula, and the third to the Khorat area. The reports are concise and contain useful geological summaries (Lee, W.M. 1923a, b, c). These reports form the important foundations of Thai geology. Lee subsequently summarized the reports and emphasized the oil potential of the country (Lee 1927).

In 1927–9 Wilhelm Credner, from Kiel University, visited many regions of Thailand that had not previously been described, and he published the most comprehensive description of the geology yet written (Credner 1935).

In 1934 the Ministry of Defence engaged two Swiss geologists to study the possibility of petroleum in the north, including the oil shales at Mae Sor and long-known oil seeps at Maung Fang. Accounts of the petrography and structure of the region were published (Heim and Hirschi 1939; Hirschi 1938, 1939; Hirschi and Heim 1938).

In 1941, a geological survey division was established in the Royal Department of Mines, but it was not until the end of the war in 1946 that the division could be staffed and a beginning made on a long-term geological programme. The Royal Department of Mines then asked the United States Government to assist in a reconnaissance study of the mineral deposits of the country. The survey was carried out between 1949 and 1950 and the result was the first authoritative account of the geology and mineral deposits of Thailand (Brown et al. 1951). Since 1951, geological mapping, research, and publication by the Department of Mineral Resources have flourished. The work has been carried

out both by Thai geologists and in collaborative projects with the German and British geological surveys. The publications are too numerous to enumerate; and lists may be found in Workman (1978), Nutalaya *et al.* (1978) and Sethaput (1956).

However, even now, Brown *et al.* (1951) remains the only comprehensive account of the country. A modern definitive geology of Thailand has yet to appear, but Bunopas (1981) gives a useful recent summary.

1.2.2. Myanmar

In 1862 Pegu, Arakan, and Tenasserim were amalgamated to form British Burma, and the whole of Burma became incorporated in the British Indian Empire in 1886. Before the association of Burma with India there were a number of geological reports on the mineral deposits and fossil localities. Many of the reports appeared in the *Journal of the Asiatic Society of Bengal*. However, it was during the Indian administration that systematic surveys and published accounts flourished. These are to be found in the memoirs and records of the Geological Survey of India. The earlier records are too numerous to enumerate here, and the reader is referred to Goossens (1978a, b) for a bibliographic summary. The early work was excellently summarized in outstanding accounts by La Touche (1913) and later by Pascoe (1950, 1959, 1964). Sir Edwin Pascoe's massive 2130-page work was published after his death in 1949. He started writing in 1933 and submitted the complete manuscript in 1939. Most of it was already set in type, but at the beginning of the war in 1939 the 2.5 tons of type were broken up and remelted for munitions. After his death, the manuscript of Volume 4, which consisted of general and geographical indexes, was unfortunately never traced. This massive work is therefore not indexed, but it represents the outstanding introduction to the pre-1933 knowledge of Myanmar.

In 1923 the University of Rangoon set up a new department of geography and geology. Its staff included such famous names as L. Dudley Stamp and H. L. Chhibber. Both were active in the field and made the results of their work widely available; for example Stamp (1927). Chhibber (1934a, b) made a landmark in the geological literature by publishing his two books, one on the geology and the other on the mineral deposits of Myanmar.

Myanmar was administratively separated from India in 1937, and eventually gained independence from Britain in 1948. Since Chhibber's work, many Burmese and other geologists have contributed significantly to our understanding of the country. The

publications are summarized by Goossens (1978). A notable summary of the geology and mineral deposits was published by Robertson Research (1975, 1977) at a time when it appeared that the government was opening up the country to international mining companies. Unfortunately this did not take place, and several recent surveys, maps, and reports prepared by overseas investigators have been suppressed. A geology of Burma has been recently published by Bender (1983).

1.2.3. Indochina (Cambodia, Laos, Vietnam)

The various provinces were unified as French Indochina in 1887, but French influence extends back to 1787. The early literature on Indochina was in French. There was only occasional research up to the end of the nineteenth century. Some of it, however, like the study of the rich flora in the coal basins of North Vietnam, was of great importance (Counillon 1914, Zeiller 1902). Before 1880, the French colonial administration showed little interest in geological research and mining development. Thus, it was that the first general work on the geology of the region (Petiton 1895) was not published until some twenty years after its completion.

The Service Géologique de L'Indochine was founded in Hanoi in 1898. Field mapping and palaeontological studies progressed rapidly, but for many years they were restricted to selected areas, most of the effort being centred on Yunnan and southern China. From 1925 onwards, emphasis was placed on the compilation of a geological map of the whole of Indochina, completed only in 1963. The publications up to 1950 are largely the fruits of the Service Géologique de L'Indochine. Individual works are too numerous to enumerate, but mention should be made of the important summaries by Fromaget (1941) and Saurin (1935, 1944). A detailed bibliography of this and the later period was compiled by Fontaine (1973). French Indochina split up into North Vietnam, South Vietnam, Laos, and Cambodia in 1956, and the newly independent countries formed their independent geological surveys. Since then, North and South Vietnam have unified and Cambodia was renamed Kampuchea before again becoming Cambodia.

The geology of North Vietnam has been extensively re-studied, with assistance from the USSR. To the extent that wartime conditions allowed, South Vietnam and Cambodia continued to perform new mapping, but little new work has been possible in Laos, and little is known still of the geology of that country. A booklet on the geology of North Vietnam was published by Tran *et al.* (1979), but nothing comparable has come recently from the other countries which formerly

comprised Indochina. Fontaine and Workman (1978) provide the most comprehensive bibliography and most recent summary of the geology and mineral resources of these countries.

1.2.4. Indonesia

A very impressive geological programme of mapping, research, and publication flourished under Dutch control, which gradually spread throughout Indonesia. The last territory to be annexed was the northern part of Sumatra as late as 1907. Many significant surveys and publications were made in the latter years of the nineteenth century, but the most important works were published in the first two decades of the twentieth century. Many eminent geologists either worked in Indonesia or took part in well-organized expeditions there. The Geological Survey of the Netherlands Indies lasted from 1850 to 1950, with its headquarters in Bandung and the Bureau of Mines in Jakarta (then Batavia). During that time, the regular Survey publication was the series *Jaarboek van het Mijnwezen*, published in Batavia. In addition, several books and many articles were published in Europe on the geology of Indonesia. Virtually everything ceased in 1941 with the outbreak of war.

Many famous Dutch geologists wrote about Indonesia, and it would be impossible to do justice to all of them here. I will mention only those authors who synthesized the earlier work and built upon it in their books. The earliest and widest compilation was by Brouwer (1925). Rutten (1927, 1932) gave a series of lectures and his books brought the attention of the world to this fascinating region of South-east Asia. Umbgrove (1949) also did much to summarize the prominent features of Indonesia. However, it was by the impressive work of Van Bemmelen (1970) that the geology of Indonesia and South-east Asia became well known. The first edition of his outstanding book was published in 1949. Van Bemmelen had been a member of the Geological Survey of the Netherlands Indies since 1927, and was working on the manuscript in Bandung in 1941 when the Japanese invaded. He was interned during the war. His manuscript could not be retrieved, and the book was begun afresh after the war. Mention should also be made of the novel gravity measurements made by Vening Meinesz in the submarine *K XIII* in 1927, 1929, and 1930. Through his work, the unique nature of the deep-sea trenches of the Indonesian region was brought to the attention of the world (Vening Meinesz 1954).

The literature of the Indonesian region is rich and varied, but for most people Van Bemmelen (1970)

will serve as a summary. Since the war and the independence of Indonesia, the Geological Survey of Indonesia has made great progress in better understanding this vast and complicated country. The book by Hamilton (1979), which summarizes most of the later work, is now widely known, and it interprets the region in modern plate tectonics terminology.

1.2.5. Philippines

The Philippines was held by Spain until 1898 and geological records of that era are generally now only of historic record. Enrique Abella-y-Casariego, the last chief of the Spanish Mining Bureau, was the ablest of the early investigators. His publications, now hard to find, are listed in Smith (1924).

In 1898 the Philippines was ceded to the United States of America. Soon after this a Mining Bureau was established. A considerable amount of information was published by the United States government, and it summarized all the useful earlier knowledge. Notably, Becker (1901) summarized all that was known to that date. The publication of Smith (1924) was a landmark in Philippine geology. He had been with the Mining Bureau since 1906, becoming chief of the division of mines in 1920. This outstanding book contains a complete bibliography of previous work of the Spanish era. Wallace E. Pratt, also of the Bureau, published extensively; for example, Pratt (1916). Filipinos began publishing; a notable example was Leopoldo A. Faustino, who joined the division of mines around 1920 and has a long list of publications (for example Faustino 1934).

The Philippine Commonwealth was established in 1934, which planned for and led to the independence of the Philippine Republic in 1946. However, no significant geological work was done around that time.

The next landmark was the publication of Corby *et al.* (1951), which contains a comprehensive account of the geology, emphasizing the oil potential. There have since been numerous publications on the geology of the Philippines, too numerous to mention, but summarized by Teves (1953, 1957), and by Aquino and Santos (1971). The most recent landmark is the publication by the Bureau of Mines (1981), summarizing the geology in book form. Volume 2 of this work is on the mineral deposits, published later (Bureau of Mines, 1986).

1.2.6. Peninsular Malaysia and Singapore

Peninsular Malaysia, formerly known as the Federation of Malaya, was under British rule until 1957. With independence it became Malaysia, and eventu-

ally included Sarawak and Sabah (North Borneo), which are jointly now referred to as East Malaysia. Singapore withdrew from Malaysia and is now an independent country.

The earliest and most comprehensive records of the geology of Peninsular Malaysia and Singapore were by Logan (1848). Between 1870 and 1903 Europeans began participating in the mining industry of the peninsula and several authors wrote on the tin deposits. However, until 1903 any geological work had been of a short-term and haphazard nature. It was against this background of the important tin-mining industry that the government appointed its first geologist, J. B. Scrivenor, who set up his office in Batu Gajah, in the foremost tin-mining valley of Kinta. He arrived in 1903 and devoted most of his working life to Malaysian geology. Until his retirement in 1930, most of the geological publications were written by Scrivenor, who laid the foundations of Malaysian geology. His other earlier works were admirably summarized in his two books (Scrivenor 1928, 1931). In 1913, W. R. Jones joined Scrivenor as a geological assistant. He studied the great Kinta Valley tin field and summarized its geology in Jones (1925). The early geological literature of Peninsular Malaysia and Singapore has been compiled by Gobbett (1968).

The Federation of Malaya Geological Survey continued to grow, and the headquarters moved to Ipoh, where the main laboratories are still sited. With independence in 1957, the Geological Survey of Malaysia gradually phased out its expatriate staff. The Survey headquarters moved to Kuala Lumpur, and assumed control of geological work in both Peninsular and East Malaysia. In addition to a large number memoirs and map bulletins, the Survey continues to map systematically and in detail the geology of the whole of Malaysia.

Geological work in Singapore comes under the Public Works Department, but work is usually limited to site investigations, and no systematic geological mapping is undertaken.

Establishment of a geology department in the University of Malaya in 1956 was an important landmark in the development of the geological sciences in the region. The professors – C. S. Pichamuthu, T. H. F. Klompé, N. S. Haile, K. F. G. Hosking, C. S. Hutchison, and P. H. Stauffer – have all made significant contributions to Asian geology. The Geological Society of Malaysia was founded in 1967 by the academic staff of the department, notably by D. J. Gobbett, and it continues to flourish.

As a successor to Scrivenor (1928, 1931), a book was published under the editorship of Gobbett and

Hutchison (1973). It remains the authoritative summary of the geology of Peninsular Malaysia and Singapore. A revised edition is currently under preparation. The Republic of Singapore (1976) has published a summary of its own geology.

1.2.7. Sarawak, Sabah, and Brunei

Accounts of geological explorations in Borneo have been given by Hatton (1886), Posewitz (1892), Schmidt (1904), and Rutter (1922). The early systematic geological studies of Sarawak, Brunei, and Sabah were made primarily by oil companies and also by companies exploring for mineral deposits. But no comprehensive accounts came until a geological survey was established.

The Geological Survey Department, British Territories in Borneo, was first established in 1949, with offices in Kuching and Kota Kinabalu, then known as Jesselton. An important landmark was the publication of Reinhard and Wenk (1951). The Shell Oil Company employed these two Swiss geologists to compile a comprehensive account of the geology of North Borneo (Sabah). The field work was completed before 1942 and the results published by the Geological Survey Department in 1951.

Shell also collaborated with the Geological Survey Department in making available the only regional compilation of the whole territory, including Sarawak, Brunei and West Sabah (Liechti *et al.* 1960). Since then a large number of reconnaissance memoirs and bulletins have been produced by the Department, and later by the Geological Survey of Malaysia. Independent Brunei is dominated by the Shell Oil Company and its geology has recently been summarized in book form by James (1984). The latest book on North-west Borneo, Sarawak, Brunei and Sabah, is by Hutchison (2005).

1.3. THE MODERN ERA OF GEOLOGICAL CO-OPERATION

Rapid advances in the geological knowledge of South-east Asia have been made since the early 1970s, as a result of planned international co-operation between oceanographers and land-based geologists within a programme of transect studies across peninsular South-east Asia and its island arcs. This programme was carried out within the International Decade of Ocean Exploration, and co-ordinated by the Committee for Co-ordination of Joint Prospecting for Mineral Resources in Asia Offshore Areas (CCOP) and the Intergovernmental Oceanographic Commission of

UNESCO (IOC) (CCOP-IOC 1974, 1980). From this programme, there has been a rapid flow of papers in international journals and books over the past decade, which has greatly improved our understanding of South-east Asia. Active offshore exploration by petroleum companies also greatly increased our knowledge of the Tertiary basins. The release of some of their data helped, for example, in the compilation of Hamilton (1979). An unnecessary reluctance by national petroleum corporations to make available regional data, has, however, greatly hindered tectonic analysts. Summaries of the Tertiary oil basin geology have been published, for example, by ASCOPE (1981). Oil companies of the region continue to release permitted information through publications of the Indonesian Petroleum Association, Jakarta, the Southeast Asia Petroleum Exploration Society, Singapore, and various regional geological societies, such as the Geological Society of Malaysia, Kuala Lumpur, which holds an annual petroleum geology seminar and publishes the papers in its bulletin series.

In 1972 the Geological Society of Malaysia began an important tradition of holding a regional conference on the geology and mineral resources of South-east Asia (GEOSEA). The proceedings have been published by various societies in the countries of the region: GEOSEA 1 (Tan 1973); GEOSEA I1 (Wiriyosujono and Sudradfat 1978); GEOSEA I11

(Nutalaya 1978); GEOSEA IV (Philippines 1981); and GEOSEA V (Malaysia 1984). The full proceedings of GEOSEA V were published in 1986 and 1988 as bulletins 19 (Teh and Paramanathan, 1986) and 20 of the Geological Society of Malaysia. GEOSEA VI was held in Jakarta in July 1987. The ninth GEOSEA conference proceedings were published as Bulletin 43 of the Geological Society of Malaysia.

Volume 1 of the *Journal of Southeast Asian Earth Sciences*, published by the Pergamon Press in Oxford and edited by B. K. Tan of the University of Malaya, appeared in 1986. This journal has been superseded by the *Journal of Asian Earth Sciences*, now published by Elsevier.

The research programmes of the International Decade of Ocean Exploration Studies in East Asian tectonics and resources (SEATAR) have resulted in several special publications, for example Barber and Wiriyosulono (1981), Hayes (1980, 1983) and Hamilton (1979). Important new work is in progress in the Philippines, the Timor area, and the South China Sea. Not all SEATAR transects have been published but those across the Banda Sea and Java to Sarawak have been published (Hutchison, 1991a, b).

Three outstanding quality maps of the region have been compiled by Gatinsky (1983), by Ray (1982) and by Pubellier et al. (2005), and the oil and gas map by ESCAP (1985).

Late Mesozoic and Cainozoic tectonic features

2.1. GENERAL

The region is covered extensively by seas (Fig. 2.1). However, not all are of similar character. The 200-m water depth is a useful demarcation isobath between shallow seas that cover the continental shelf and the deeper more oceanic seas. South-east Asia has a number of marginal seas (Bostrom, 1978), where the water depth exceeds 3 km; these abyssal plains are floored by oceanic crust.

The marginal seas are separated by island or remnant arcs (Karig, 1972) made up either of volcanic or accretionary wedge material. In addition, the arcs may be complicated by the inclusion of continental fragments that have been rifted from the larger continental masses of Asia and Australia, and carried away by seafloor spreading.

The two main shelf sea regions are the Sunda Shelf, which is a submarine extension of continental South-east Asia, connecting the peninsular landmasses with Sumatra and Borneo, and the northwards extension of Australia as the Sahul Shelf, extending north to the Timor Trough. This shelf extends to link up with New Guinea via the Arafura Sea. Detached fragments of the Australian platform occur in the islands of Sumba and eastern Sulawesi (Pigram and Panggabean 1984; Hutchison 1987b). Continental fragments occupy much of the South China Sea and are assumed to have been detached from the China shelf between Hainan and Taiwan. These fragments include north-east Palawan, and several submarine banks such as the Reed Bank and Luconia Shoals (Taylor and Hayes 1983).

Low sea levels during the Pleistocene Great Ice Age resulted in the landmasses extending 21 thousand years ago to the 120 metre isobath (Hanebuth et al. 2000; Voris, 2000). The continental shelf, extending to the 200 metre isobath is probably wholly a result of the low sea levels and shoreface sedimentation during the past 140 thousands of years (Fig. 2.1).

2.2. WALLACE'S LINE AND ITS MEANING

The eminent English naturalist Alfred Russel Wallace divided South-east Asia into Oriental and Australian

faunal domains by a line which became known as Wallace's Line (Fig. 2.29). West of the line, the fauna and flora have distinct Asian affinities: east and south-east of the line they have Australian or Oceanic affinities. From 1863 to 1880 Wallace drew the line between Bali and Lombok, extending between Borneo and Sulawesi (then called Celebes) and between the Philippines and Indonesia. In 1910, he revised it to lie east of Sulawesi (George 1982). Sulawesi presented difficulties to Wallace, which have even today not been resolved, for its fauna and flora seem to have affinities with both Asia and Australia. The present-day consensus is that western Sulawesi was formerly a part of South-east Asia, and the eastern part, which is largely of ophiolite, was uplifted as a result of collision of the Banggai-Sula Spur, a westwards displaced continental fragment of New Guinea. Wallace's Line, in geological terms, therefore runs through central Sulawesi (Audley-Charles 1982).

All terrains east of Wallace's Line were previously remote from South-east Asia, and have been drifting northwards since about 90 Ma when Australia separated from Antarctica and began its northwards drift which accelerated to about 8 cm a⁻¹. Their flora and fauna had ancestral roots in the supercontinent of Gondwanaland. The collisions of these terrains with those of South-east Asia began only about 15 Ma, and the Sahul Shelf collided with the Timor Trough at only 3 Ma (Audley-Charles 1982). As the Australian terrains approached South-east Asia, dispersal of fauna and flora across the narrowing seas that separated them caused some intermingling of the previously distinct fauna and flora. This explains the difficulties that Wallace had in drawing a definite line in the neighbourhood of Sulawesi (Fig. 2.29).

The picture is further complicated because India separated from Gondwanaland (Australia) about 120 Ma, and began its drift northwards, eventually to collide with Tibet about 50 Ma. This allowed fauna and flora of Gondwanaland origins to reach Asia and disperse into South-east Asia. There is also a possibility that other smaller landmasses of South-east Asia may have drifted away from Gondwanaland. There is a growing

Cainozoic sedimentary basins

It is desirable to have some system of classification applicable to the thick accumulations of predominantly Tertiary strata, which have become the focus of attention for oil, gas, and coal exploration and exploitation in the region. Any mechanism which can produce a depression in the earth's crust is a potential basin-forming agent. Obviously, faulting is the main mechanism, but it is important to differentiate the kinds of faults and to consider their origin in relation to the tectonic evolution of the region.

3.1. BASIN CLASSIFICATION

Two attempts at classification of South-east Asian basins appeared independently in 1975 (Murphy 1975; Soeparjadi *et al.* 1975). The more detailed one by Murphy (1975) subdivides the basins into 'shelfal', 'continental margin', 'archipelagic', and 'marginal seas', but some of these are too broad to be useful. Soeparjadi *et al.* (1975) used such categories as 'outer arc', 'foreland', 'interior cratonic', and 'open shelf on continental margin'. Some of these terms are open to misinterpretation. 'Interior cratonic' rightly applies to such basins as the Sichuan 'Red' basin of the Yangzi platform and to the Khorat-Vientiane basin of Indosinia, but should not be applied to the Borneo region.

It is preferable to use terminologies related to the elements of plate tectonics. The general tectonics-based classifications of Bally and Snelson (1980), Horn (1980) and Kingston *et al.* (1983) are useful, and may be followed where appropriate. These two classifications are broadly similar despite different terminologies. One important difference is in regard to the mechanisms of rifting. Kingston *et al.* (1983) interpret the various back-arc basins of convergent plate margins as resulting from wrench faulting incorporating a component of divergence. Thus, they refer to the back-arc basins of Sumatra and the Malay basin as 'wrench or shear basins'. This is in agreement with the analysis of Wood (1985) and has been followed by Hutchison (1985, 1986a) and displayed in Figs. 3.1 and 3.2. We know that these back-arc

basins have both extensional and wrench histories (Crostellla 1981). The detailed code of Kingston *et al.* (1983) is difficult to apply without extensive knowledge of any particular basin, so that of Bally and Snelson (1980) is more readily applicable. They are, however, broadly complementary.

Hutchison (1985) showed that several basins peripheral to Borneo have been complicated by collision between microcontinents. He therefore created an additional category: 'basins on or peripheral to continental fragments'. This category remains nevertheless somewhat obscure.

A fundamental problem is whether the basins which lie behind the volcanic arc are genetically related to the arc-trench system. The 'back arc' category may be nothing more than a geographical description. From the analysis of Wood (1985) and Tapponnier *et al.* (1982), these basins are now known to have formed by wrench tectonics and widespread rifting of the South China Sea. The same problem applies to the classification of 'marginal seas'. Few, if any represent 'back-arc' ocean floor spreading (Hutchison 1986a). Many represent fragments of larger oceans trapped behind younger arc-trench systems. Others, like the South China Sea Basin, result from rifting of a continental shelf (Taylor and Hayes 1983). The use of a term to classify a basin does not, therefore, mean that we understand its origin. In many cases the terms are geographically descriptive and unrelated to genesis.

The locations of the major Tertiary basins are shown in Fig. 3.1. Their names, suggested classification, and average geothermal gradients are shown in Fig. 3.2 and listed in Table 3.1.

The fitting of geological details into various favoured models (Table 3.1) has been contrived and attempts at relating the stress fields of the faulting patterns (e.g. Wood, 1985) and basin shapes to the active plate margins have been largely unsuccessful. The fore-arc category alone remains viable, but only where the plate margin has not been reorganized throughout the basin evolution, e.g. Sibolga, Bengkalis and South Java. All others, i.e. the 'back-arc' basins, remain genetically enigmatic and defy meaningful plate tectonic classification.

Phanerozoic tectonic framework

4.1. INTRODUCTION

South-east Asia represents the ultimate eastern extension of the Palaeo-Tethys (Şengör 1984; Şengör and Hsü 1984; Belov *et al.* 1986). The elimination of the Palaeo-Tethys resulted in the formation of the Late Triassic–Early Jurassic Indosinian (Cimmeride) Orogeny. The Neo-Tethys Ocean has not been totally eliminated by the Palaeogene Alpide (Himalayan) orogeny and is partly extant in the Indian Ocean. Only to the north and east of India has it been converted into suture zones in Tibet and Myanmar (Fig. 4.1).

The Cainozoic fault motions, discussed in Chapter 2, have largely moulded the present geological framework and made more difficult the unravelling of the earlier geological history and plate motions (Stauffer 1974, 1983).

The South-east Asian or *Indosinian orogenic system* (Indochinese Cimmerides of Şengör, 1984) is the least well known of the Cimmeride orogenic system and the gaps in our knowledge are attributed to many reasons, predominantly the equatorial climate leading to deep weathering and lush tropical rainforest, complex political boundaries, and a general lack of access. Nevertheless, its geological framework has emerged, though important essential details are lacking, particularly in the structural aspects. The Indosinian orogen shows great similarities to the western extensions of the Cimmeride orogenic system. The differences are unique to South-east Asia, where the Palaeo-Tethyan Ocean was widest and hence a greater separation between Gondwanaland and Asia existed.

4.2. ANALYSIS OF THE TECTONIC FRAMEWORK

Many attempts have been made to analyse the tectonic evolution of South-east Asia. The following are

considered to be the most significant: Hutchison (1973c, 1987a), Stauffer (1974, 1983, 1985), Acharyya (1978b), Gatinsky *et al.* (1978, 1984), Şengör (1979, 1984), Ridd (1980), Bunopas (1981), Gatinsky (1981, 1986a, b), Mitchell (1981), Şengör and Hsü (1984), Metcalfe (1986), and Gatinsky and Hutchison (1987). Throughout this series of papers, there has developed a growing awareness that South-east Asia is composed of a number of relatively rigid Precambrian continental blocks, marginally and extensively overlain by Phanerozoic platform strata, and separated by highly deformed mobile belts composed of strongly compressed deep water strata deposited upon oceanic crust of the Palaeo-Tethys or its branches. Commonly, they are closely associated with ophiolite and mélange belts marking suture zones (Figs 4.1 and 4.2). Subduction activity within the oceans and along their margins with the continental blocks is documented by magmatic arcs; their nature and ages are summarized in Fig. 4.2. The simplified tectonic map of Şengör and Hsü is shown here as Fig. 4.1. It shows that the region represents an assemblage of terrains ('tectonic collage' of Şengör 1984¹), welded together along narrow suture zones, and finally consolidated into the Cimmeride orogenic system in Jurassic times.

A plate boundary has long been recognized extending southwards from Yunnan as the Changning–Shuangjiang Suture (Huang *et al.* 1984), through Thailand and the Malay Peninsula. Hutchison (1975a) named it the *Uttaradit–Luang Prabang Line* in Thailand and Laos, and it extends northwards to the vicinity of Dien Bien Phu, where it appears to be represented by a major right-lateral strike-slip fault. In Peninsular Malaysia he named it the *Bentong–Raub Line* (Fig. 4.2). Unfortunately, Şengör (1984), and Şengör and Hsü (1984) did not maintain the terminology of Hutchison (1975a); their alternative nomenclature is shown in Fig. 4.1.

From the tectono-stratigraphic entity distributions, it may be concluded that this boundary was of major

1. 'Collage' is a term used exclusively for an artistic artifact, and is hardly appropriate. 'Terrain assemblage' is preferable, and means the same.

Terrains of Cathaysian affinity

The blocks of Cathaysian affinity shown in Fig. 4.2 are: South China (6), Indosinia (7), the small microcontinent of Phu Hoat (8), the West Borneo Basement (11), and Eastern Malaya (Eastmal) (12). Sumatra (10) has mixed affinities and is included for convenience wholly in this chapter.

5.1. SOUTH CHINA BLOCK

The South China Block is composite, subdivided into the Yangzi Platform on the west, succeeded eastwards by the Nan-ling 'Caledonian' fold belt and then by the eastern Min-Zhe volcanic province. The major structural features of the block are illustrated in Fig. 5.1, summarized from Yang *et al.* (1986), and in the diagrammatic cross-sections (Fig. 5.2), modified after Terman *et al.* (1974). Although the Precambrian is extensively covered by Infracambrian and Phanerozoic formations, it is well exposed in the Yangzi Gorges, in the uplifted margins of the platform, and in the massifs of North Vietnam.

5.1.1. Yangzi Platform basement

South China is dominated geologically by the large Yangzi Platform built upon a stable predominantly Proterozoic basement, bordered on the south-west by the Trungson mobile belt, and on the south-east by the Nan-ling or Xiang-Gui 'Caledonian' fold belt. The Platform is separated on the north from the North China Block by the Qin-Ling suture and fold belt, and bordered on the west by the Longmen Shan (Fig. 5.1). The geology has been summarized by Yang *et al.* (1986).

The basement of the Yangzi Platform is heterogeneous and was consolidated by the Jinningian Orogeny (1050–850 Ma). The basement is known to be as old as Archean in the Dabie Uplift and Early Proterozoic beneath the Sichuan Basin, represented by outcrops of the Kongling Group in the Yangzi Gorge. The most mobile parts of the basement occur along the south-east margin in the Jiangnan Uplift, chiefly of Upper Proterozoic rocks. The western margin of the

platform is exposed as the Kham–Yunnan Uplift, formed of Lower to Middle Proterozoic rocks (Fig. 5.2).

5.1.1.1. Archean of the Dabie Uplift

The Dabie Group is of at least 15 km thickness of amphibolite facies metamorphosed basic volcanics and sedimentary rocks including banded iron formations. The oldest radiometric date is 3120 Ma and others of 2500 and 2080 indicate Proterozoic metamorphic events. The Group is unconformably overlain by the Proterozoic Hongan Group in Hubei.

5.1.1.2. Proterozoic basement of South China

In South China, Lower Proterozoic rocks are known only from the Yangzi Platform basement (Yang *et al.* 1986; Wang and Qiao 1984). The high grade Dabie Group may include both Archean and Lower Proterozoic. It is overlain unconformably by the Susong Group, which contains marble and phosphate rock, with radiometric dates of 1850–2343 Ma.

In the interior of the Platform, the Yangzi Gorges expose the Sandouping (Kongling) Group which is directly overlain by the Sinian System. The Sandouping Group is of gneiss and amphibolite followed by schist and marble, then again by gneiss and amphibolite. Total thickness is 3 km and the metamorphic age has been dated at 1688 Ma.

In the Kham–Yunnan Uplift, the Dahongshan Group lies unconformably beneath the Middle Proterozoic Kunyang Group. It is exposed at Xinping near the Red River fault zone, and characterized by meta-sediments and alkaline volcanic rocks. Radiometric ages range from 1704 to 1958 Ma.

The Hekou Formation of north-east Yunnan and south-east Sichuan is of similar lithology. It is overlain by the Middle Proterozoic Huili Group.

Middle to Upper Proterozoic rocks occur widely in South China. They pass up into the non-metamorphic Infracambrian Sinian System, which forms the lower part of the platform cover of the basement.

The Sandouping Group of the Dabie Uplift is unconformably overlain by the 6 km thick non-metamorphic Shennongjia Group, predominantly of

Gondwana affinity terrains

6.1. INTRODUCTION

The western part of continental South-east Asia is composed of terrains which have distinct Gondwana¹ affinities in their Upper Carboniferous and Permian successions. Such terrains are characterized ideally by:

1. Carbo-Permian diamictites (tilloids);
2. Late Permian *Glossopteris* and related flora
3. Cool or cold water Late Carboniferous to Early Permian fauna, such as the Late Carboniferous *Stepanoviella* and the Early Permian *Lytvolasma*, *Taeniothaerus*, *Monodioxodina*, and *Uraloceras*. These are in strong contrast to the Lower Permian warm water *Neoschwagerina*, *Dasycladaceae*, *Codiaceae*, and *Gymnocodiaceae* (Wang and Mu 1983).

All three together are taken to infer a definite attachment to Gondwanaland during pre-Carboniferous times. The exact time of separation of the lithospheric blocks from the Carbo-Permian glaciated continent may have been as long ago as Late Carboniferous, but the latest detachment was that of India in Early Cretaceous time, 130 Ma (Veevers 1984). Mixed affinities of diamictites, *Glossopteris* and warm water fauna may indicate that the lithospheric block was removed sufficiently far from Gondwanaland by Early Permian time so that the shelf seas became warm. Mixed affinities in Middle Eastern countries is possible because of the narrow width of the Palaeo-Tethys Ocean, across which fauna and flora could easily have dispersed. But this was impossible in South-east Asia where the Palaeo-Tethys was wide.

Some terrains contain apparently mixed Gondwana and Cathaysia flora, for example in Tibet (Li *et al.* 1984), Turkey and Saudi Arabia (Şengör 1985a), and it is not yet certain how to interpret them. The western Tethys around Turkey and Saudi Arabia can

be explained by the narrow Tethys with little climate contrast across the Ocean. It may also be that there has been misidentification of fragmentary fossil leaf collections. This problem needs further research.

The several lithospheric blocks are not all defined to be of Gondwana ancestry with equal certainty, but two out of the three criteria listed above should be taken as fairly certain indications of the affinity.

The rifting from Gondwanaland is assumed to have been of Atlantic-type rifted passive margins, so that the Gondwana terrains are characterized by miogeoclinal or platform Palaeozoic successions, in places known to overlie Precambrian basement, but commonly the basement is covered and unknown in outcrop. There is a remarkable stratigraphic similarity between the terrains, from Tibet in the north, to Sinoburmalya² in the south.

The lithospheric blocks are separated by suture zones one from another. The terrains are described in order from the north to south of the region.

6.2. THE TIBET PLATEAU

The Tibet or Qinghai-Xizang Plateau forms its own terrain, referred to as the Tethys-Himalayan tectonic domain (Huang *et al.* 1984). It is wholly of Gondwana affinity and is sutured along its northern border along the Longmuco-Yushu suture onto the Cathaysian terrain of Hoh Xil-Bayan Har, which forms the southern province of the East Kunlun region. The contact zone then bends southerly towards the east where it is known as the Jinsha Jiang-Teng Tiaohe suture, which separates the Gondwana terrain on the west, here known as the Yunlijig-Wuliang Shan, from the Cathaysian Yangzi Paraplatform on the east (Fig. 6.1).

The Tibet Plateau is separated on its southern margin from the Himalayan province by the major

1. Named in 1872 by H. B. Mendlicott, director of the Geological Survey, after *Gond*, a central India aboriginal hill tribe and *Wana*, Sanskrit for forest. Thus, *Gondwanaland*, the supercontinent of which India was part more than 130 Ma.

2. Named by Gatinsky *et al.* (1984) for an elongated terrain extending from Yunnan, through Myanmar to the Malay Peninsula. It is equivalent to *Sibumasu* of Metcalfe (1984a, 1986) and *Shan-Thai* of Bunopas (1981). Metcalfe's *Sibumasu* specifically includes north-west Sumatra.

Ophiolites and sutures

There are numerous linear or arcuate belts of ophiolitic and mélangé rocks in South-east Asia. They have been reviewed by Hutchison (1975a), and Wiryosujono and Tjokrosapoetro (1978). The geographic distribution of the main belts is shown in Fig. 7.1, and each is described in turn. They are divided into two categories.

- Ophiolites occurring in suture zones within the Eurasian continent, which appear to have resulted from the closure of former ocean basins.
- Ophiolites which occur along continental margins with oceans and marginal seas, which have not yet formed suture zones. The northwards motion of Australia continues to push the Indonesian–Philippine archipelago towards continental China, so that these ophiolitic belts will form future suture zones in the orogenic welt between Australia and China (Hamilton 1979).

Many ophiolites exist solely as blocks forming an integral part of a mélangé formation. The subduction mélangé wedge is of deformed sheared and faulted sediments, and ophiolitic blocks overlying the shallow parts of a Benioff zone. Although Hamilton (1979) holds that mélangé wedges result from gravitational flow trench-wards from the outer arc ridge, the mélanges of Nias and other accretionary prisms are related to palaeo highs. Moore and Karig (1980) concluded that thrust-related shearing on the lowermost trench slope is the most likely mechanism for creating mélanges, and this thrusting facilitates the upwards movement of ophiolites when the sediment cover is still thin.

Some large ophiolitic masses have very well preserved stratigraphy extending from oceanic upper mantle, through gabbro layer, basaltic layer, to pelagic sediments, and even into fore-arc sedimentary sequences. They must be regarded as intact up-faulted or obducted masses of marginal basin lithosphere. There is therefore a spectrum from large coherent masses to the smallest clasts in a tectonic or sedimentary mélangé.

The processes for coherent uplift or obduction of ophiolite have been modelled by many, including

Gass (1977), and Searle and Stevens (1984). Collision of a continental massif such as Australia with an island arc such as Indonesia, followed by partial underthrusting of the depressed miogeoclinal leading edge of the continent, would provide an attractive mechanism for uplifting oceanic lithosphere. The continental edge cannot continue to subduct, and the volcanic arc is quickly extinguished, as illustrated by the Alor–Wetar–Romang sector north of Timor. The plate margin may continue to be tectonically active, but with a changed pattern. This model has been used for the Papuan ophiolite (Davies 1981; Davies and Jaques 1984), and it may prove attractive elsewhere. It is certainly logical to conclude that an uplifted dense ophiolitic mass needs to be isostatically supported by low density continental rocks which have been underthrust beneath it.

7.1. OPHIOLITIC SUTURES

7.1.1. Donqiao ophiolite of the Bangong–Dengqen–Nujiang suture

This suture is located about 300 km north of the Indus–Yarlung–Zangbo suture (Fig. 6.1). South of the suture, between Donqiao and Gyanco, the ophiolite outcrops are closely spaced, but spread over a width of 70 km. The array represents a single flat-lying ophiolite nappe obducted at least 180 km southwards from the suture (Chang *et al.* 1986). A cross-section across the suture zone shows, from bottom to top, a succession of Jurassic flysch and ophiolitic and Palaeozoic metasedimentary rocks (Fig. 7.2). They all stand in flat tectonic contact with each other and are overlain unconformably by Albian–Aptian redbeds and volcanic rocks (Girardeau *et al.* 1984).

The fossiliferous Jurassic flysch has undergone low-grade (sericite-chlorite) metamorphism and is folded round sub-horizontal axes and has developed a slaty S_1 cleavage. A sub-vertical late crenulation S_2 cleavage is superimposed and related to late symmetrical folds.

The ophiolite everywhere tectonically overlies the Jurassic flysch. Its basal contact is marked by tectonic

The great Sunda–Pacific volcanic arcs

One of the world's great Cainozoic volcano-plutonic arcs characterizes the convergent plate margin of the Indian Plate against Eurasia, and the marginal seas that characterize the South-west Pacific. Together, they form a nearly continuous arc from the Andaman Sea, along the margins of Sundaland, to the northern Australian margin, then north-wards through the Philippines to Okinawa.

On the Indian side, the arc has resulted from subduction of the India–Indian Ocean–Australia Plate. Before continental India began its collision with Eurasia to form the Himalayas, there was pre-Eocene subduction of oceanic lithosphere beneath Tibet to form the Gandise arc, which abruptly terminated with the arrival of the Indian Platform at the subduction zone. The Burman Volcanic Arc was active until subduction ceased along the Indo-Burman Ranges former trench location. Southwards from Myanmar, oceanic lithosphere of the Indian Ocean continues to subduct beneath Sundaland (Fig. 2.21), and the Indonesian Volcanic Arc is continuous from Sumatra, through Java to Flores. North of Timor, there is an extinct and uplifted sector of the arc from Alor to Romang, because of the arrival of continental Australia at the subduction zone, but activity continues eastwards in the Banda Arc, though the volcanic rocks show geochemical evidence of under-thrust Australia. Most of this eastern sector will progressively extinguish as Australia pushes northwards.

On the Pacific side, the great volcanic girdle continues northwards from Halmahera and Sangihe through the Philippines, where the arcs are related to subduction of the marginal sea lithosphere of the Celebes, West Philippine Basin, and South China Sea Basin. Opposing directed subductions are in the process of active reorganization, and the volcanic arcs will change in time and place. The Taiwan sector is extinct because of the underthrust continental shelf of China, but eastwards the West Philippine Basin lithosphere subducts at the Ryukyu Trench to give the active Okinawa volcanic arc.

8.1. GANDISE ARC OF THE INDUS–ZANGBO–YARLUNG SYSTEM

The volcano-plutonic arc occupies southern Tibet along the Gandise–Nyainqentangtha Mountain Range (Fig. 9.1). Its eastern end turns south-east at Chayu, and in the west it extends to Ladakh in Kashmir, being 1000 km east–west and about 100 km north–south (Geological Bureau of Xizang 1982).

The arc plutonic part comprises the huge Gandise intermediate to acid intrusive complex (Chapter 9), and the volcanic part comprises intermediate-acid volcanic rocks of the Upper Jurassic–Lower Cretaceous Sangri Group and the widespread Cretaceous–Tertiary Linzizong Formation volcanic series. The plutonic and volcanic rocks are closely associated, comprising a range from augite diorite, granodiorite and granitoids, and andesite, dacite, rhyolite, and trachyte.

The *Sangri Group* volcanic rocks are associated with a sequence of Late Jurassic to Early Cretaceous neritic carbonates and clastic sediments. It is chiefly composed of andesite, dacite, keratophyre, and a huge amount of andesitic and dacitic crystal tuff, volcanic breccia, and other pyroclastic varieties. The chemistry suggests an island-arc calc-alkaline series.

The Linzizong Formation unconformably overlies a sequence of folded Late Cretaceous red clastic rocks. Below the volcanic rocks is 50–60 m thickness of conglomerate on the Qinghai–Tibet Highway, 40 km north of Lhasa (Geological Bureau of Xizang 1982). The *Linzizong* Formation is largely of lava and tuff, intercalated with continental clastic strata. The *Linzizong Volcanic Series* is over 2500 m thick, composed mainly of andesite-dacite-rhyolite lavas and pyroclastic rocks, lying unconformably on shale of the Late Cretaceous Takena Formation (Wang 1984).

There are three major cycles, each one dominated by andesite and andesitic pyroclastic rocks in the lower part, and by dacite and rhyolitic pyroclastic rocks, and tuffaceous sandstone and conglomerate in the upper part. Cycle I is up to 966 m thick. The lower part is of a spectrum from tholeiitic, through calc-

Granite and associated plutonic rocks

South-east Asia has an unusually strong concentration of granite batholiths, bearing witness to the various orogenic phases which have moulded the geology of the region. The granitoids are regarded as rift-, subduction-, or collision-related (Mitchell 1977, 1979). All types are important in the region, and the granites have been responsible for much of the mineral wealth of the region (Hosking 1977). In particular the tin and tungsten deposits, for which the region is renowned, are genetically related to Mesozoic S-type granites, both of the Late Triassic–Early Jurassic Indosinian Orogeny, and of the Late Yenshanian (Cretaceous) earth movements.

A general review of the granites of South-east Asia is to be found in Cobbing *et al.* (1992), Schwartz *et al.* (1995) and Hutchison (1983b), and this is expanded in this chapter to include the Himalayas and South China. The subduction-related granitoid terrains are characterized by an extended calc-alkaline series including gabbro, diorite, granodiorite, and granite of I-type; whereas the collision-related batholiths are restricted to S-type granite (Hutchison 1982c). The Late Yenshanian granites of South China and the South China Sea region are rift-related, but are distinctly not of the A-type of White and Chappell (1983); they are indistinguishable from S-type and constitute a hitherto unrecognized category of sub-alkaline crustally-derived rift-related granite of great importance to mineralization (Hutchison 1987c). The distribution of magnetite and ilmenite series granitoids is discussed by Takahashi *et al.* (1980).

9.1. TIBET

The granitoids of southern Tibet and Nepal form distinct belts parallel to the Indus–Yarlung–Zangbo suture (Fig. 9.1): the Gandise batholith belt north of the suture, and the Lhagoi–Kangri and Himalayan belts south of the suture (Liu 1984; Debon *et al.* 1986).

9.1.1. Gandise or Transhimalayan batholith system

This is the major magmatic arc of the region. It extends from Swat (Pakistan) through Ladakh and

Kailas along the Gandise Range of the Lhasa Block (Valdiya 1984). It has a range of ages. The oldest group, from 82 to 113 Ma, has a mean value of about 93 Ma (Late Cretaceous). It characterizes both the north and south Zangbo units (Debon *et al.* 1986). An early basic pluton of the southern edge of the Gandise belt gave $^{39}\text{Ar}/^{40}\text{Ar}$ ages of 90–110 Ma (Albian to Cenomanian), whereas lavas from the upper unconformable series are clearly younger, 60 Ma, or Palaeocene (Maluski *et al.* 1984). The youngest group varies from 40 to 60 Ma and averages about 50 Ma (Eocene). It includes all the plutonic units in a 90 km section from Yangbajain to south of Quxu. The dating is a summary of a variety of methods and is very likely to correspond with the emplacement ages. A very young Rb–Sr date of 17 Ma is thought to be anomalous (Debon *et al.* 1984).

The initial $^{87}\text{Sr}/^{86}\text{Sr}$ ratios are characteristically low, within the range 0.7033–0.7036 (Valdaya 1984), but Jin and Xu (1984) gave a higher range of 0.7047–0.7057 for Tibet, now extended to 0.7036–0.7071, with an average of 0.705 (Debon *et al.* 1986). On the whole, the northern plutons have higher initial ratios than the main Gandise batholith.

The Gandise batholith system is composite and exhibits a spectrum, including in decreasing order of importance, adamellite, quartz monzodiorite, granodiorite, and to a lesser extent gabbro, quartz diorite, granite, quartz monzo-diorite, and monzonite. Magmatic differentiation, characterized by two superimposed stages of evolution, and hybridization processes involving both basic and acid magmas, can account for the genesis of the different plutonic units (Debon *et al.* 1986). The rocks are principally composed of microcline, plagioclase, biotite with $\text{Mg}/(\text{Mg} + \text{Fe})$ from 0.45 to 0.61, hornblende, magnetite, sphene, and rare ilmenite. Orthopyroxene and clinopyroxene are restricted to gabbro. The whole batholith has I-type characteristics (Liu 1984; Jin and Xu 1984; Debon *et al.* 1986). From west to east and from south to north, the system becomes younger and more acidic. Granodiorite is predominant west and granite predominant east of Qushul, while diorite commonly

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